

## Earth and Planetary Materials

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Spring 2013

Lecture 5  
2013.01.23

## Simple structures – AX

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**CN = 8: the CsCl structure**

**CN = 6: the NaCl structure**

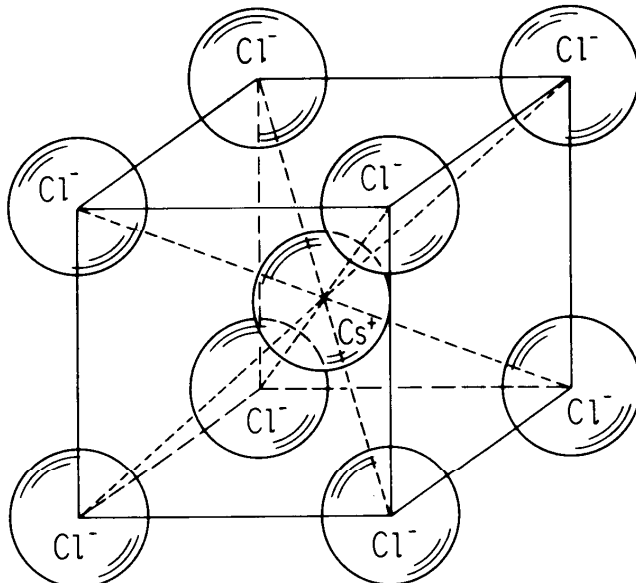
**CN = 4: the wurtzite and sphalerite structures**

## AX structure (CN = 8)

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### CN = 8: the CsCl structure

- Based on a **simple cubic (SC)** arrangement of anions
- All the CN=8 (cubic) sites are filled
- The Cs(Cl)8 cubes share faces



**Figure 2-3** CsCl structure. This structure is made up of two interpenetrating PC lattices. The unit cell of CsCl may be shown with either Cs<sup>+</sup> or Cl<sup>-</sup> in the PC positions. All ions are in cubic coordination.

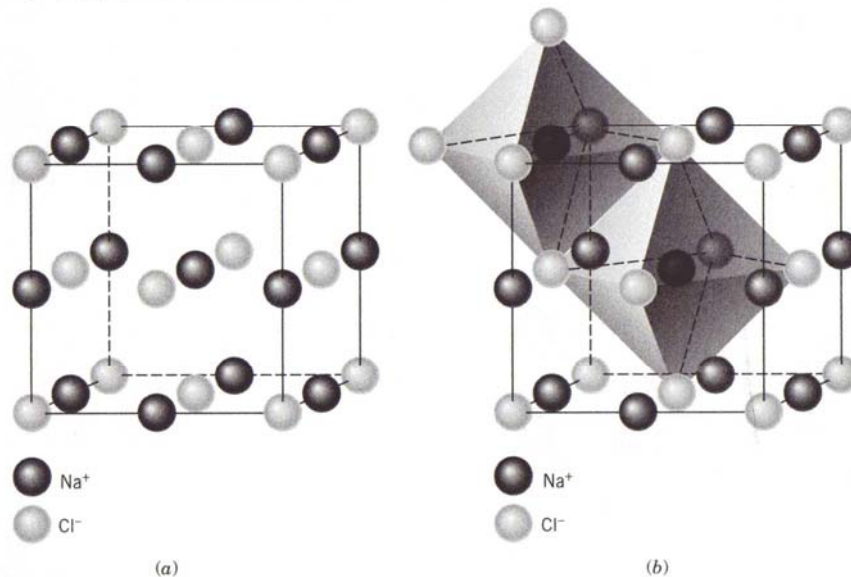
## AX structure (CN = 6)

### **CN = 6: the NaCl structure**

(halite structure, rock salt structure)

- Based on a CCP arrangement of anions, which corresponds to an FCC unit cell
- The Na(Cl)6 octahedra share edges

**FIG. 4.17** (a) The structure of NaCl, halite. The  $\text{Na}^+$  and  $\text{Cl}^-$  ions are arranged in a face-centered cubic lattice. (b) The same structure showing the edge-sharing octahedra about the  $\text{Na}^+$ . Similar edge-sharing octahedra could be drawn about  $\text{Cl}^-$ .



## AX structure (CN = 6) – common minerals

- alkali halides with radius ratios up to RbCl
- alkaline earth oxides, including MgO (periclase)
- Some transition metal oxides, including FeO (Wüstite)

1 H 1.008																	2 He 4.003														
3 Li 6.941	IIA										5 B 10.81	6 C 12.01	7 N 14.01	8 O 16.00	9 F 19.00	10 Ne 20.18															
11 Na 22.99	12 Mg 24.31											13 Al 26.98	14 Si 28.09	15 P 30.97	16 S 32.07	17 Cl 35.45	18 Ar 39.95														
19 K 39.10	20 Ca 40.08	21 Sc 44.96	22 Ti 47.87	23 V 50.94	24 Cr 52.00	25 Mn 54.94	26 Fe 55.85	27 Co 58.93	28 Ni 58.69	29 Cu 63.55	30 Zn 65.39	31 Ga 69.72	32 Ge 72.61	33 As 74.92	34 Se 78.96	35 Br 79.90	36 Kr 83.80														
37 Rb 85.47	38 Sr 87.62	39 Y 88.91	40 Zr 91.22	41 Nb 92.91	42 Mo 95.94	43 Tc (99)	44 Ru 101.1	45 Rh 102.9	46 Pd 106.4	47 Ag 107.9	48 Cd 112.4	49 In 114.8	50 Sn 118.7	51 Sb 121.8	52 Te 127.6	53 I 126.9	54 Xe 131.3														
55 Cs 132.9	56 Ba 137.3	57 La 138.9	58 Ce 140.1	59 Pr 140.9	60 Nd 144.2	61 Pm (145)	62 Sm 150.4	63 Eu 151.9	64 Gd 157.3	65 Tb 158.9	66 Dy 162.5	67 Ho 164.9	68 Er 167.3	69 Tm 168.9	70 Yb 173.0	71 Lu 174.9	72 Hf 178.5	73 Ta 180.9	74 W 183.8	75 Re 186.2	76 Os 190.2	77 Ir 192.2	78 Pt 195.1	79 Au 197.0	80 Hg 200.6	81 Tl 204.4	82 Pb 207.2	83 Bi 209.0	84 Po (209)	85 At (210)	86 Rn (222)
87 Fr (223)	88 Ra (226)	89 Ac (227)	90 Th (232)	91 Pa (231)	92 U (238)	93 Np (237)	94 Pu (244)	95 Am (243)	96 Cm (247)	97 Bk (247)	98 Cf (251)	99 Es (252)	100 Fm (257)	101 Md (258)	102 No (259)	103 Lr (260)	104 Rf (261)	105 Db (262)	106 Sg (266)	107 Bh (264)	108 Hs (269)	109 Mt (268)	110 Uun (271)	111 Uuu (272)	112 Uub (277)						

## AX structure (CN = 4)

### CN = 4: the wurtzite and sphalerite structures

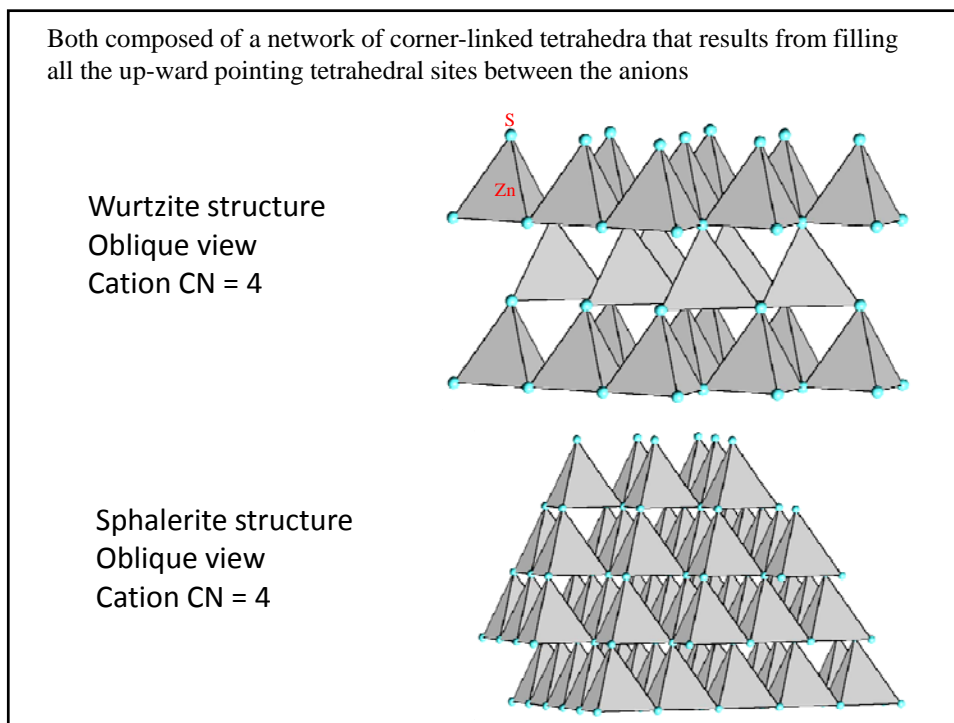
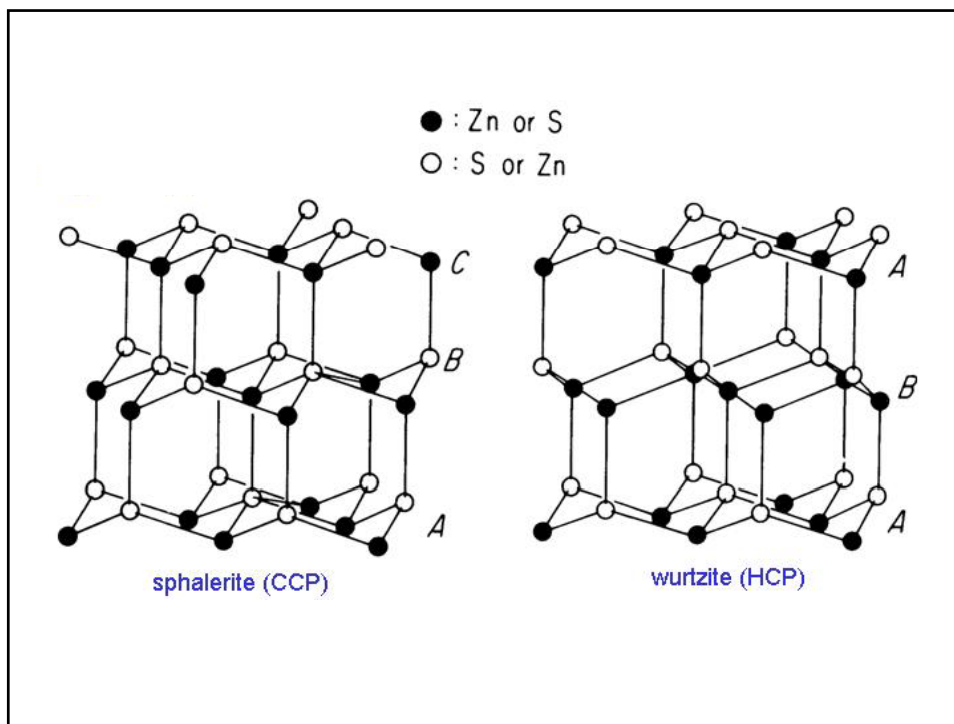
- polymorphs of ZnS

wurtzite



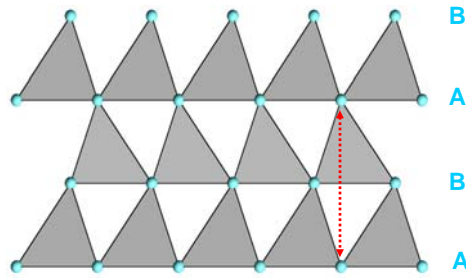
sphalerite



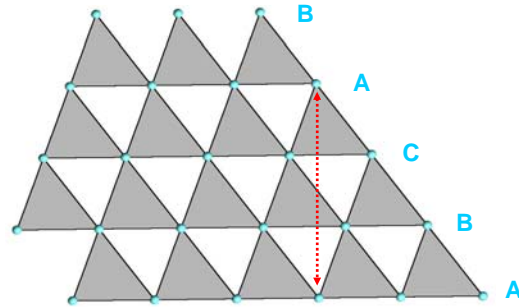


Differ in the packing of the anion ( $S^{2-}$ )

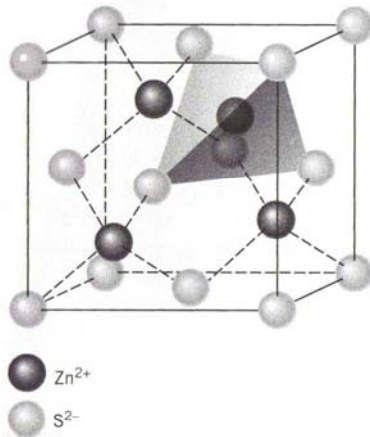
Wurtzite structure  
Side view  
HCP array of S



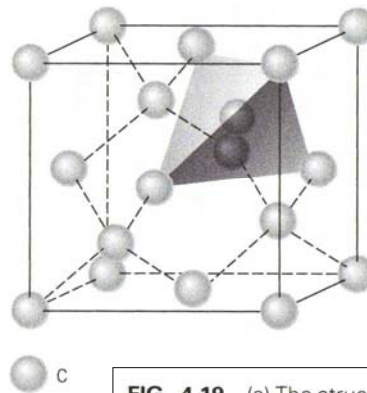
Sphalerite structure  
Side view  
CCP array of S



sphalerite ( $ZnS$ ) structure



diamond structure



**FIG. 4.19** (a) The structure of sphalerite,  $ZnS$ . Both  $Zn^{2+}$  and  $S^{2-}$  are in a face-centered cubic array.  $Zn^{2+}$  is in tetrahedral coordination with four  $S^{2-}$  neighbors (see also Fig. 4.10c). (b) The structure of diamond, C. The coordination of C by four nearest carbon neighbors is tetrahedral. (See also Fig. 3.18.)

## Simple structures – $AX_2$

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CN = 8: the fluorite structure

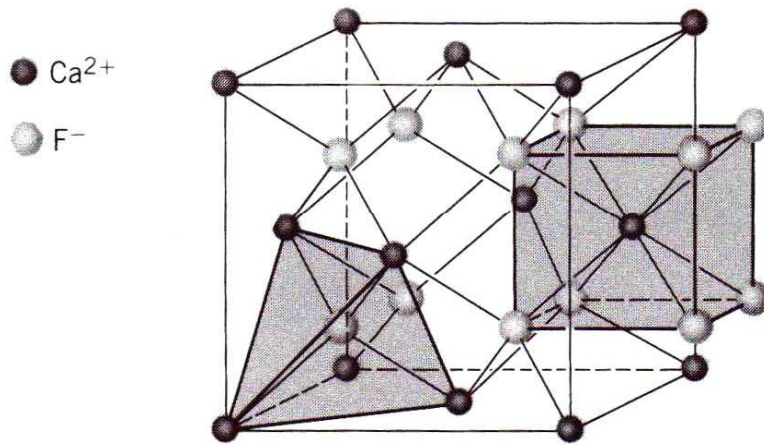
CN = 6: the rutile structure

### $AX_2$ structure (CN = 8)

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#### **CN = 8: the fluorite structure**

- anions in a simple cubic (SC) array
- cations fill one-half the cubic sites (cation CN = 8)
- anion CN = 4
- cation polyhedra (cubes) share edges



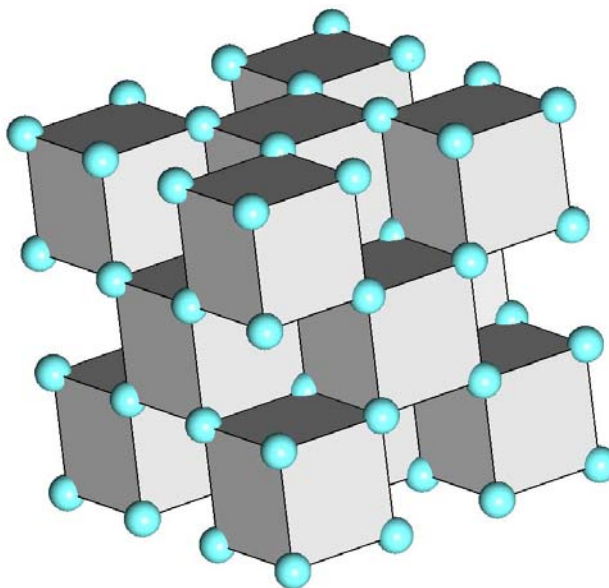
Fluorite structure

**FIG. 4.20** The structure of fluorite,  $\text{CaF}_2$ .  $\text{Ca}^{2+}$  ions are arranged in a face-centered cubic lattice. The  $\text{F}^-$  ions are in simple cubic packing (SCP) with  $\text{Ca}^{2+}$  occupying the voids at the centers of *alternating* cubic interstices.

### Polyhedral representation of the fluorite structure

Ca in FCC arrangement

CN = 8

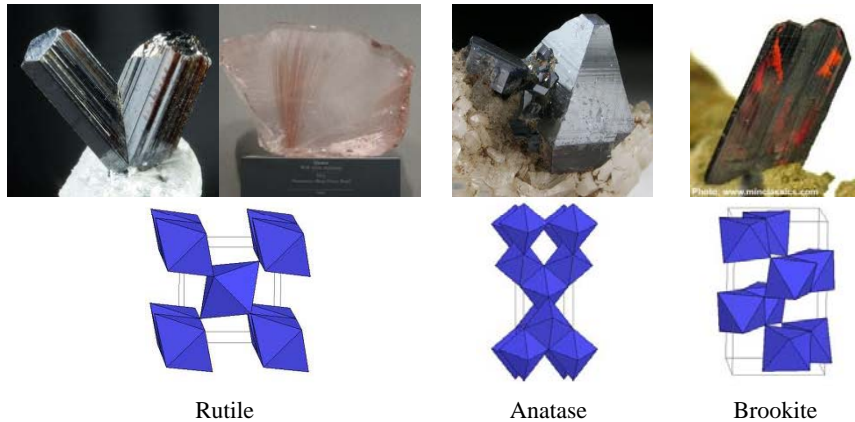




## AX<sub>2</sub> structure (CN = 6)

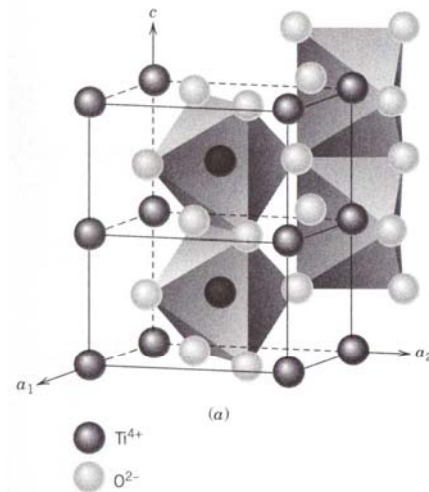
CN = 6: The rutile structure

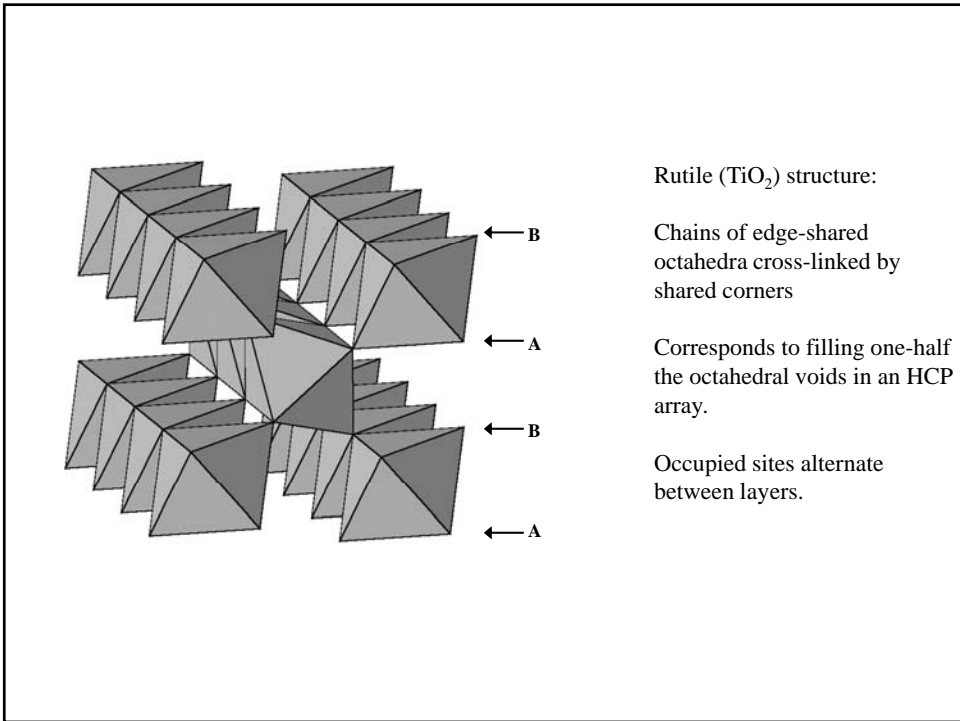
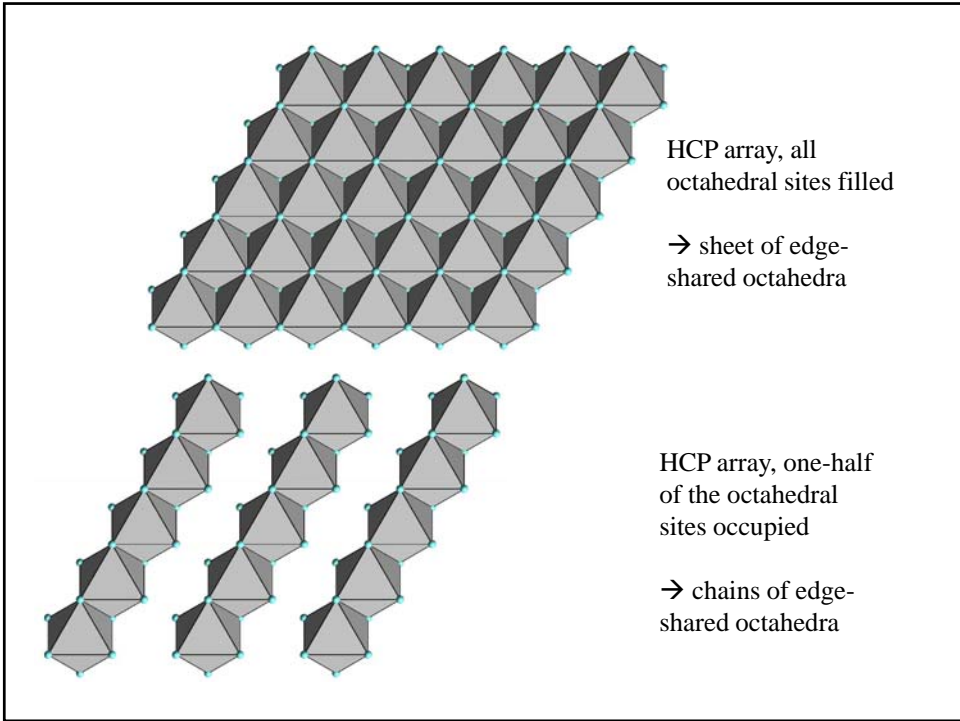
### Polymorphes of TiO<sub>2</sub>



## AX<sub>2</sub> structure (CN = 6)

- anions are in an approximately HCP array
- cations occur as chains of edge-sharing octahedra
- chains are connected via corners
- chains are formed by filling one-half the octahedral sites in a HCP sheet





## Oxides and hydroxides

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Perovskite structure ( $ABO_3$ )

Spinel structure ( $AB_2O_4$ )

Corundum group

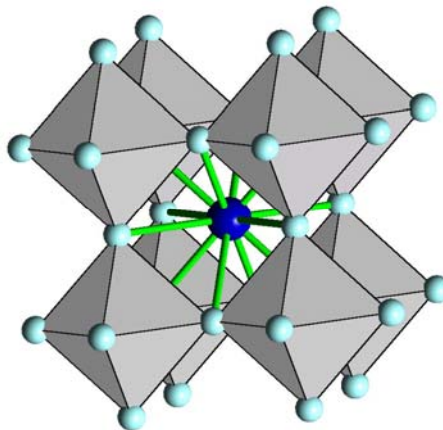
Simple layered hydroxides

Oxyhydroxides

## Perovskite structure ( $ABO_3$ )

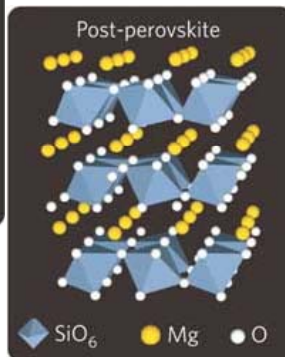
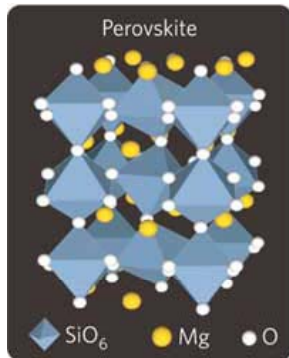
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- Named for the mineral perovskite:  $CaTiO_3$
- A cations are large, CN = 12
- B cations have CN = 6, form a network of corner-linked octahedra
- This structure is adopted by a very large number of compounds



## Perovskite structure ( $ABO_3$ )

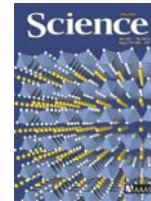
- Mg-silicate perovskite  $MgSiO_3$  – primary mineral in the upper part of lower mantle



### Post-Perovskite Phase Transition in $MgSiO_3$

Motohiko Murakami,<sup>1\*</sup> Kei Hirose,<sup>1\*</sup> Katsuyuki Kawamura,<sup>1</sup> Nagayoshi Sata,<sup>2</sup> Yasuo Ohishi<sup>3</sup>

In situ x-ray diffraction measurements of  $MgSiO_3$  were performed at high pressure and temperature similar to the conditions at Earth's core-mantle boundary. Results demonstrate that  $MgSiO_3$  perovskite transforms to a new high-pressure form with stacked  $SiO_6$ -octahedral sheet structure above 125 gigapascals and 2500 kelvin (2700-kilometer depth near the base of the mantle) with an increase in density of 1.0 to 1.2%. The origin of the  $D''$  seismic discontinuity may be attributed to this post-perovskite phase transition. The new phase may have large elastic anisotropy and develop preferred orientation with platy crystal shape in the shear flow that can cause strong seismic anisotropy below the  $D''$  discontinuity.



Murakami et al. (2004) Science, 304:855.

## Spinel structure ( $AB_2O_4$ )

- The spinel-type oxides are very common accessory (minor) minerals in igneous, metamorphic, and sedimentary rocks
- Generally dark-colored and opaque
- based on a CCP array of oxygen
- cations occur in tetrahedral and octahedral coordination
- the ratio of tetrahedra : octahedra = 1:2

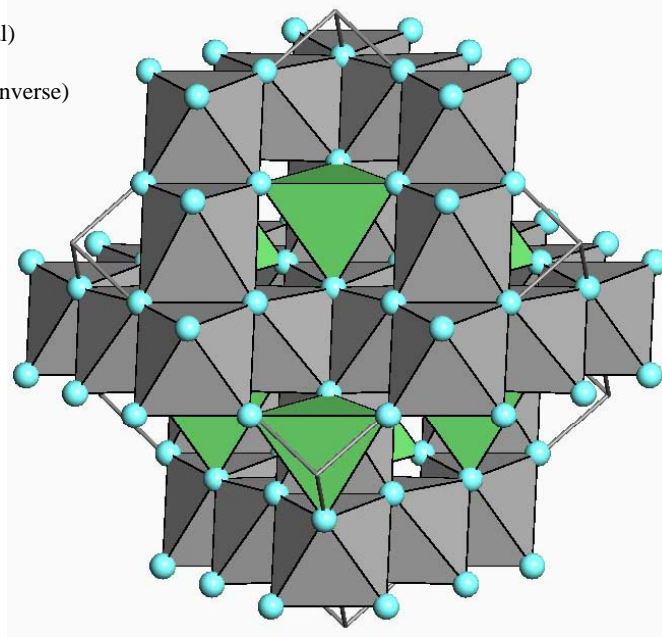
## Spinel structure ( $AB_2O_4$ )

- Chemistry of spinels is complex (Klein p372-373)

Spinel can be

1. "normal": octahedral sites filled with "B" cations
2. "inverse": "B" cations fill the tetrahedral sites, "A" and "B" cations on the octahedral sites

Spinel structure:  
 $A^{(4)}B^{(6)}_2O_4$  (normal)  
 $B^{(4)}[B^{(6)}A^{(6)}]O_4$  (inverse)



## Spinel structure ( $AB_2O_4$ )

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- Most spinels lie somewhere between these extremes
- Most also exhibit Mg/Fe solid solution  
Examples: chromite - magnesiochromite  
magnetite - magnesioferrite

## Corundum structures

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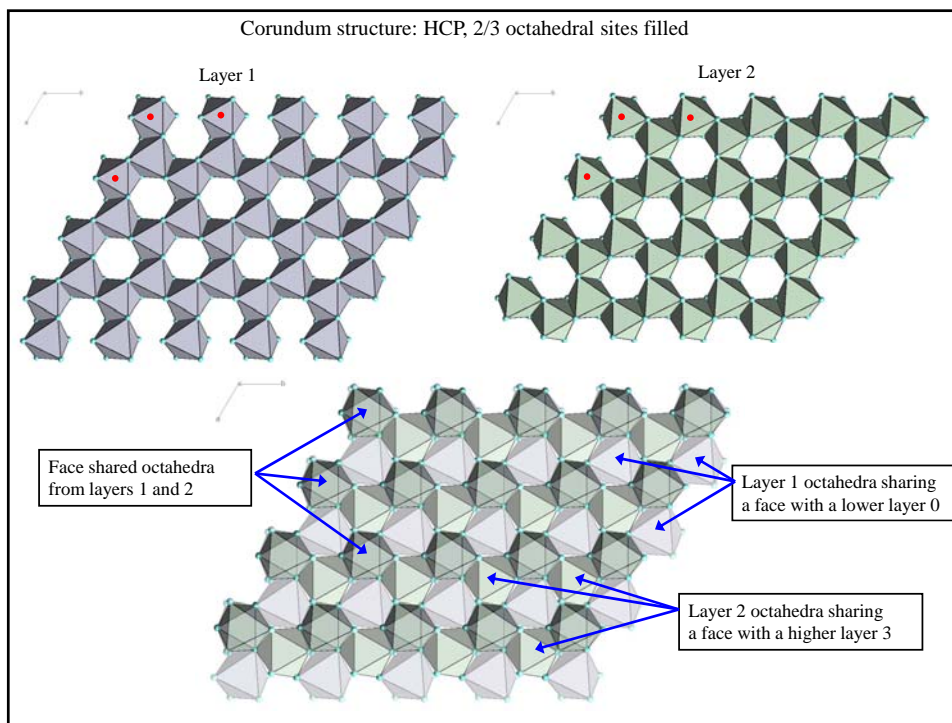
Minerals with the corundum structure include:

- corundum  $Al_2O_3$ ; varieties include  
sapphire (blue), ruby (red)
- hematite  $Fe_2O_3$
- ilmenite  $FeTiO_3$

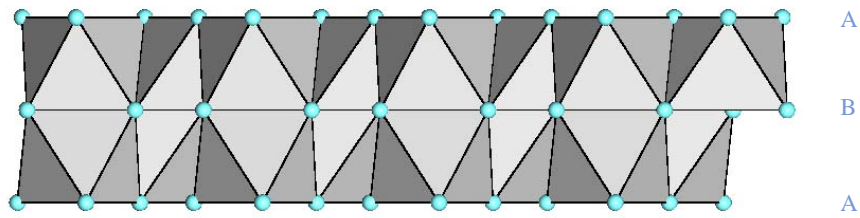


## Corundum structures

- based on an HCP array of oxygen.
- all cation have CN=6
- cations fill 2/3 of the octahedral sites
- each octahedron shares a face with one other octahedron



Corundum structure, side view: HCP array of anions



### Simple layered hydroxides

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- Common minerals in soil environments and represent important structural components of more complex minerals such as the phyllosilicates (sheet silicates).



## Simple layered hydroxides

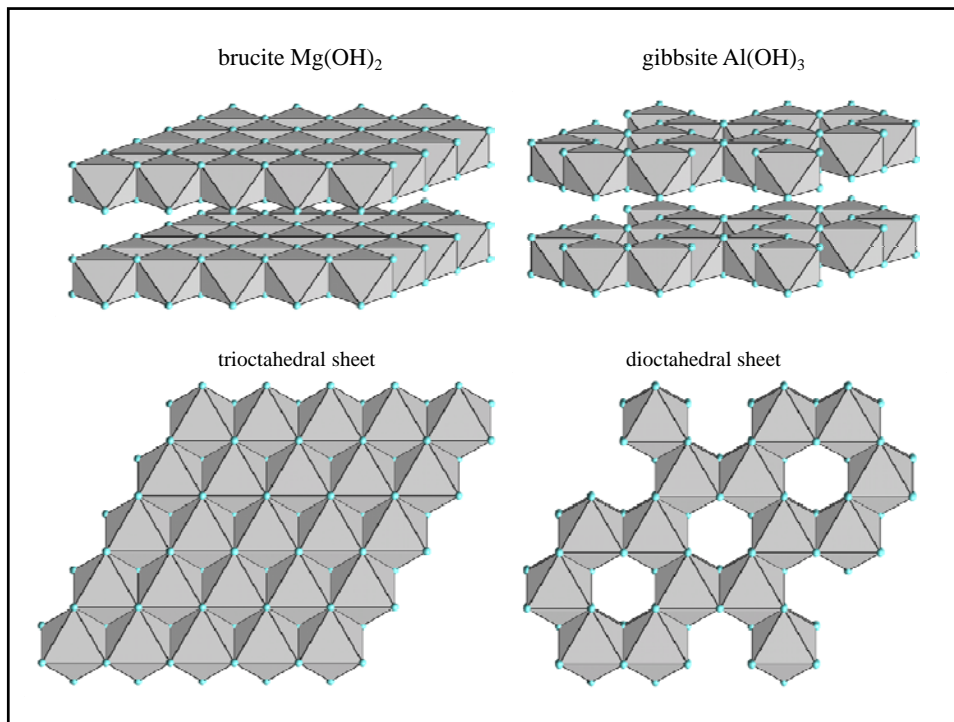
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- Based on an HCP array of hydroxyl (OH<sup>-</sup>) ions
- cations have CN=6
- Every other layer of octahedra filled
- Layers held together by weak Van der Waals and/or hydrogen bonding

## Simple layered hydroxides

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- Two main types depending on the charge of the cations
  - +2 cations fill all the octahedra in a sheet: "trioctahedral"  
Example: brucite
  - +3 cations fill 2/3 of the octahedra in a sheet: "dioctahedral"  
Example: gibbsite



## Oxyhydroxides

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- many are based on HCP anion arrays
- Typically, metal cations with CN = 6
- Examples: goethite  $FeO(OH)$  - a common soil and alteration mineral isostructural with diaspore  $AlO(OH)$  (a component of bauxite ores).

goethite FeOOH

H of hydroxyl (OH) group

