How is the ionosphere formed? a. Photoionization

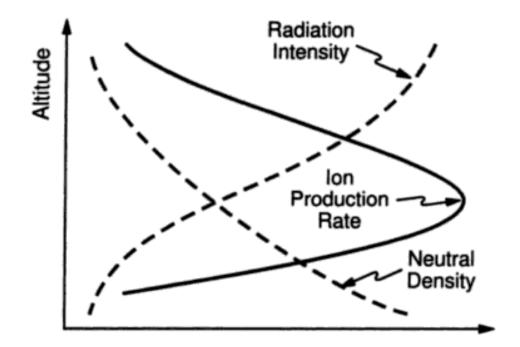
Ionization Rate = (radiation intensity) x (ionization cross section) x (density)

$$q_{z} = I_{z}\sigma_{i}n_{z}$$

$$q_{z} = I_{\infty}\sigma_{i}n_{o}\exp\left[-\frac{z-z_{o}}{H} - \frac{\sigma_{i}n_{o}H}{\mu}\exp\left[-\frac{z-z_{o}}{H}\right]\right]$$

Note: σ_i is ionization cross section and σ is absorption cross section. There is much wavelength and species dependent chemistry detail buried in σ_i

Formation of an ionized layer:

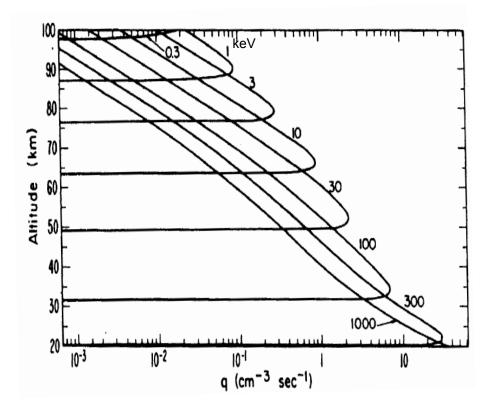


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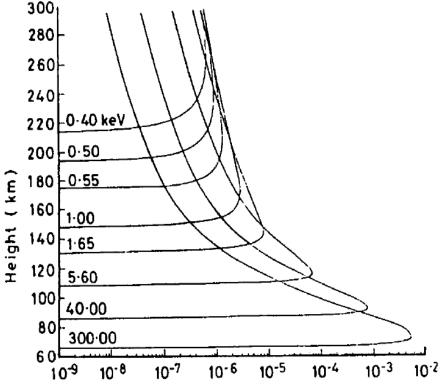
$$q_{z} = I_{\infty}\sigma_{i}n_{o}\exp\left[-\frac{z-z_{o}}{H} - \frac{\sigma n_{o}H}{\mu}\exp\left[-\frac{z-z_{o}}{H}\right]\right]$$

How is the ionosphere formed?b. Impact ionizationIonization Rate - Energy Dependent:

Due to proton precipitation



Due to electron precipitation

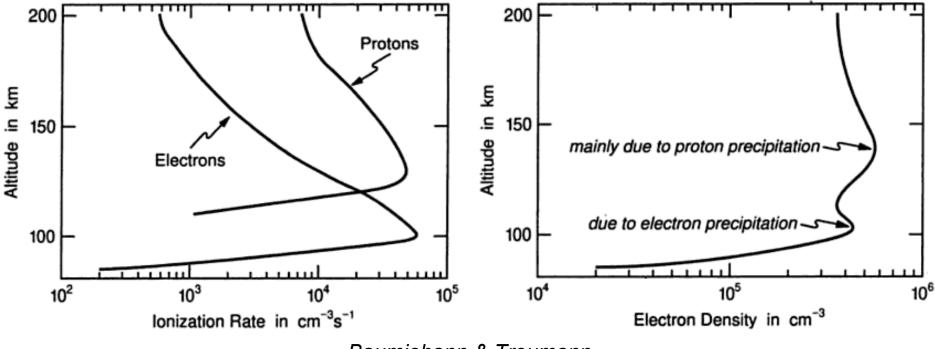


How is the ionosphere formed?

b. Impact ionization

Ionization Rate Integrated over all precipitation energies:

Total ionization due to precipitation over diffuse aurora:



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Note: These processes are most important at high latitudes (due to magnetic morphology), and at night when photoionization does not contribute!

Summing up:

Production of electrons from ionization of the neutral atmosphere from solar radiation and impacts from precipitating electrons and ions is counteracted by the processes of recombination and attachment.

$$\frac{dn_e}{dt} = q_{v,e} - \alpha_r n_e^2 - \beta_r n_e$$

Considering a system in equilibrium, i.e. $dn_e/dt = 0$ We find two simple proportionalities that determine the electron number density

Lower lonosphere

In the lower ionosphere, recombination is the dominant loss process for electrons. Hence:

$$\frac{dn_e}{dt} = 0 \quad , \quad q_{v,e} - \alpha_r n_e^2 - \beta n_e = 0$$
$$n_e = \left(\frac{q_{v,e}}{\alpha_r}\right)^{1/2}$$

At lower altitudes the electron density is proportional to the square root of the ionization rate.

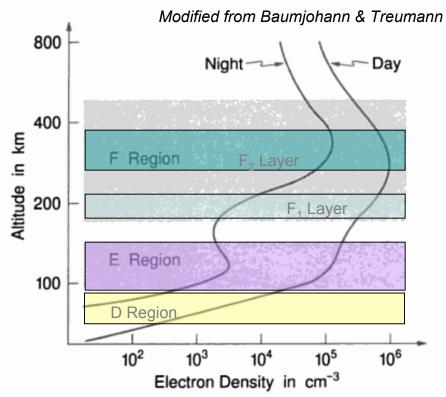
Upper Ionosphere

In the upper ionosphere, attachment is the dominant loss process for electrons. Hence:

$$\frac{dn_e}{dt} = 0 \quad , \quad q_{v,e} - \alpha \bigwedge_e^0 - \beta_r n_e = 0$$
$$n_e = \frac{q_{v,e}}{\beta_r}$$

At higher altitudes the electron density is proportional to the ionization rate.

Layers of the lonosphere:



D-region: Below 90 km, very weakly ionized and dominated by neutral gas dynamics (not a plasma)

E-region: Peaks at 110 km. Mainly produced by long wavelength UV (90nm) ionizing molecular Oxygen. At high latitudes impact ionization also plays important role. Electron density greatly reduced at night.

F-region: F_1 peaks at ~200 km, and is formed by shorter λ UV due to composition. Like the E-region, it has a strong diurnal signature. F_2 peaks at ~300 km, at high altitudes only attachment counteracts ionization (which is limited by neutral density).