Planetary Atmospheres

Structure Composition Clouds Photochemistry Meteorology Atmospheric Escape





Photochemistry

We can characterize chemical reactions in the atmosphere in the following way:

- 1. Photolysis: Molecular breakup directly driven by solar radiation (also referred to as *photodissociation*)
- 2. Photoionization: Reactions that result in the ionization of atoms and molecules
- 3. Recombination: Direct/indirect reversing of the photolysis and photoionization reactions
- 4. Dissociative Recombination: Reversing the process of photoionization via breaking a chemical bond
- 5. Charge Exchange: Direct electron exchange between a close passing ion and neutral
- 6. Atom-Ion Interchange: Interaction between an ion and atom that results in compositional alteration of the ion.

Photolysis

Oxygen in the Earth atmosphere processed by photons:

(1)
$$O_2 + hv \rightarrow O + O$$
 for $\lambda < 175$ nm
 $\frac{d[O]}{dt} = 2[O_2]J_1(z)$ Production rate of O

Where $J_i(z)$ is the reaction rate for a reaction 'i' as a function of altitude, and [atom or molecule] indicates the number per unit volume

$$J_i(z) = \int \sigma_{x_v} \mathcal{F}_v e^{-\tau_v(z)/\mu_\theta} dv$$

Since the number of photons decreases exponentially with depth penetrated into the atmosphere, production of O increases with altitude even though $[O_2]$ increases as you approach the surface.

Recombination

Direct two body recombination reverses photolysis:

$$O + O \rightarrow O_2 + hv \tag{2}$$

However, this reaction is slow, so three body processes dominate instead:

$$O + O + M \rightarrow O_2 + M$$
 (3)

$$O_2 + O + M \rightarrow O_3 + M \tag{4}$$

Where the reaction rates can be written:

$$\frac{d[O_2]}{dt} = ?$$
$$\frac{d[O_2]}{dt} = [O]^2 [M] k_{r3}$$

 k_{ri} is the reaction rate dependent on the collision rate (thus T) of the molecules

Photoionization

Oxygen and Nitrogen in the Earth atmosphere ionized by photons:

$$O_2 + h\nu \rightarrow O_2^+ + e^-$$
$$N_2 + h\nu \rightarrow N_2^+ + e^-$$
$$O + h\nu \rightarrow O^+ + e^-$$

However, these products are efficiently processed via charge exchange and atom-ion interchange to yield mostly NO⁺ and O_2^+

$$N_{2}^{+} + O_{2} \rightarrow N_{2} + O_{2}^{+}$$
$$O^{+} + O_{2} \rightarrow O_{2}^{+} + O$$
$$N_{2}^{+} + O \rightarrow NO^{+} + N$$

Ion Loss: Recombination

Dissociative Recombination:

$$O_2^+ + e^- \rightarrow O + O$$

 $NO^+ + e^- \rightarrow N + O$

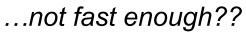
Radiative Recombination is much less efficient:

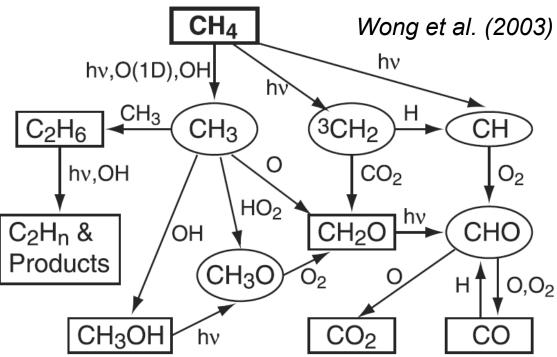
$$O^+ + e^- \rightarrow O + hv$$

Hence rapid processes like charge exchange and atom-ion interchange quickly replace the produced ions with dominant ions that can undergo dissociative recombination

Methane on Mars

Destroyed by photochemistry
 → 300–600 yr lifetime



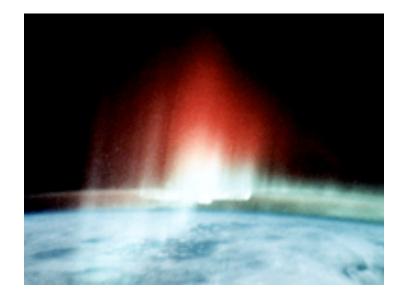


• Bar-Nun & Dimitrov (2006) argued that methane could also be produced photochemically

→ controversial

Planetary Atmospheres

Structure Composition Clouds Photochemistry Meteorology Atmospheric Escape

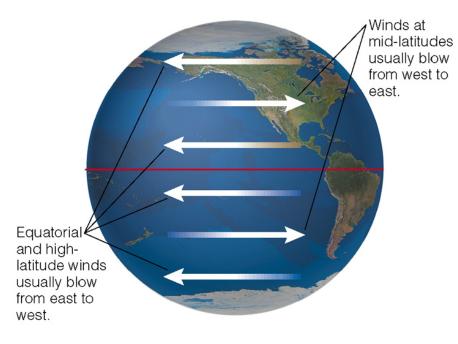




What are Weather and Climate?

weather – short-term changes in wind, clouds, temperature, and pressure in an atmosphere at a given location

climate – long-term average of the weather at a given location



- These are Earth' s **global wind** patterns or circulation
 - local weather systems move along with them
 - weather moves from W to E at mid-latitudes in N hemisphere
- Two factors cause these patterns
 - atmospheric heating
 - planetary rotation

Atmospheric Dynamics

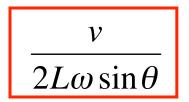
Everything Starts with the Navier-Stokes Equation:

$$\frac{D\vec{v}}{Dt} \equiv \frac{\partial\vec{v}}{\partial t} + \vec{v} \cdot \nabla\vec{v} = -\frac{1}{\rho}\nabla P + \vec{g}_p + \frac{\mu}{\rho}\nabla^2\vec{v} + f_c\vec{v} \times \hat{z}$$
Material Inertial Gravity Coriolis
Derivative Term Gravity Coriolis
$$\begin{array}{ccc}
\text{Local} & \text{Pressure} \\
\text{Derivative} & \text{Gradient} \\
\end{array}$$

$$\begin{array}{ccc}
\mu & \text{is the dynamic viscosity} \\
\frac{\mu}{\rho} = v_v & \text{is the dynamic viscosity} \\
v_v & \text{is the kinematic viscosity} \\
\end{array}$$

Atmospheric dynamics: Rotation

- Coriolis effect objects moving on a rotating planet get deflected (e.g. cyclones)
- Angular momentum as an object moves farther away from the pole, *r* increases, so to conserve angular momentum *w* decreases (it moves backwards relative to the rotation rate)
- Coriolis acceleration = $2 \omega v \sin(\theta)$ θ is latitude
- How important is the Coriolis effect?

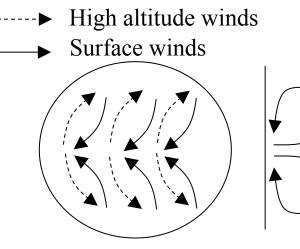


Rossby number = The ratio of inertial forces to Coriolis forces, a measure of relative importance.

e.g. Jupiter v~100 m/s, L~10,000 km we get ~1/30 so important

Coriolis Effect + Hadley Cells

Coriolis effect is complicated by fact that parcels of atmosphere rise and fall due to buoyancy creating Hadley cells (equator is hotter than the poles)

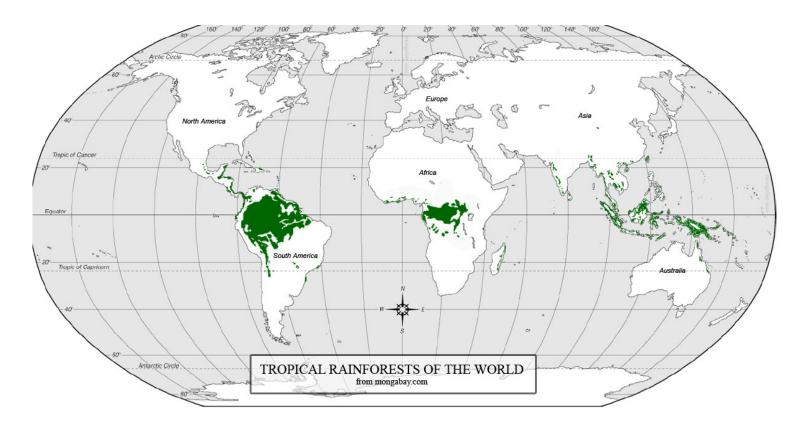


- The result is that the atmosphere is broken up into several Hadley cells (see diagram)
- How many cells depends on the Rossby number (size, rotation rate)

Slow rotator e.g. Venus Medium rotator e.g. Earth Fast rotator e.g. Jupiter Ro >>1(assumes v=100 m/s) $Ro \sim 1/4$

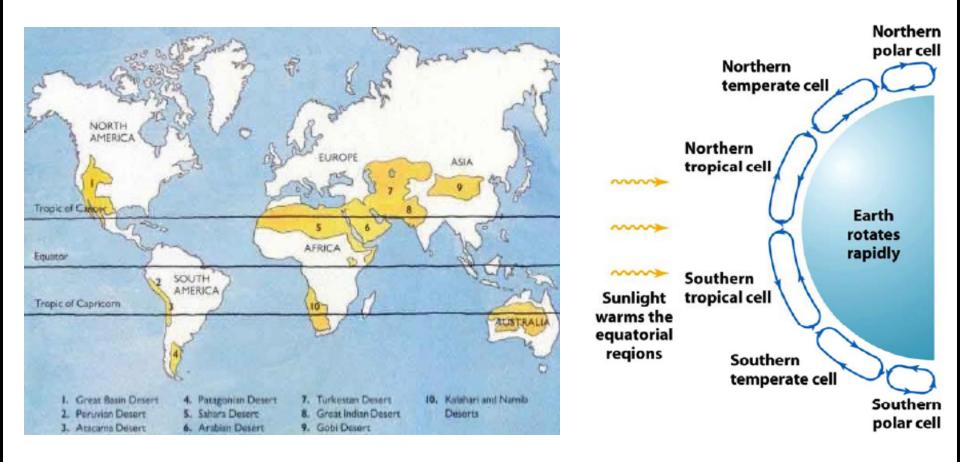
Global Wind Patterns: Hadley Cells

- Air is heated at the equator and rises
- As it rises it cools, clouds condense and generates rain



Global Wind Patterns: Hadley Cells

- Air (now dry) is pushed away from equatorial upwelling and moved toward the tropics
- As it descends, it heats up dry hotter air helps create deserts



Global Wind Patterns: Hadley Cells

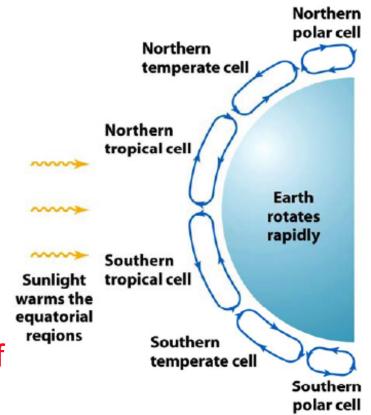
Polar cells work the same way

- Cold air descends over the poles
- Flows along the ground and eventually warms from being in contact with the surface
- Rises at ~60° N/S latitude

Ferrell Cells (Temperate Cells)

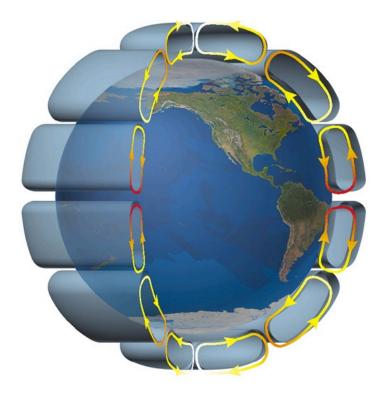
- Sandwiched between the polar and tropical Hadley cells
- Driven by their motion

The simple equator-to-poles convective motion is complicated by the Coriolis force due to the rotation of the Earth



Global Wind Patterns

- On Earth, the Coriolis effect breaks each circulation cell into three separate cells
 - winds move either W to E or E to W

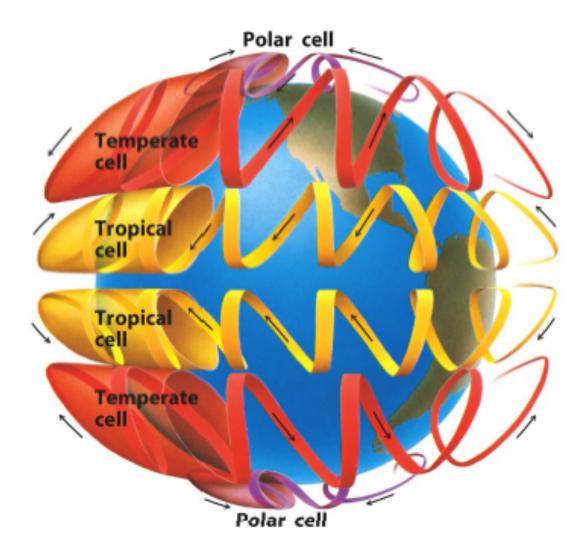


Coriolis effect not strong on Mars & Venus

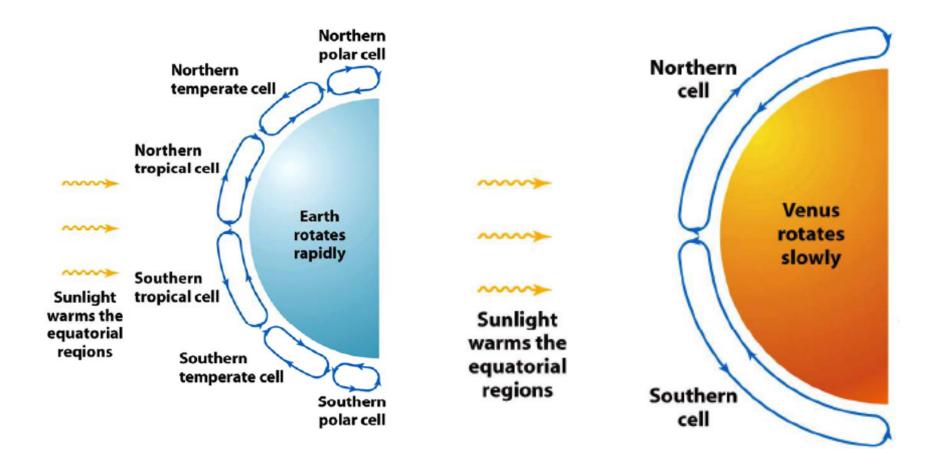
- Mars is too small
- · Venus rotates too slowly

Rapid rotators such as the giant planets are significantly affected by the Coriolis force.

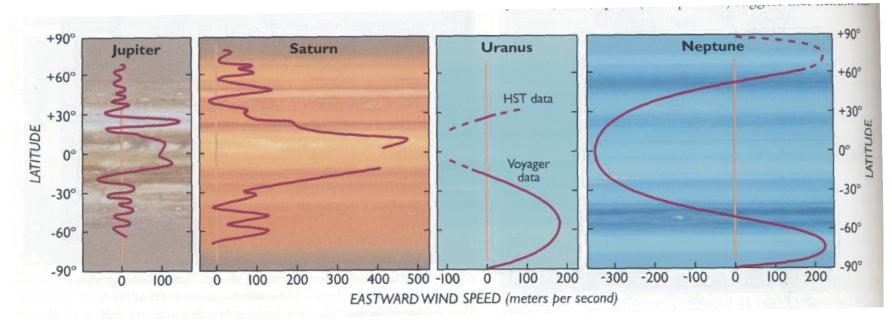
Global Wind Patterns



Comparisons



Zonal Winds



- The reason Jupiter, Saturn, Uranus and Neptune have bands is because of rapid rotations (periods ~ 10 hrs)
- The winds in each band can be measured by following individual objects (e.g. clouds)
- Winds alternate between prograde (eastwards) and retrograde (westwards)